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J. M. C. Plane. A new time-resolved model of the mesospheric Na layer: constraints on the meteor input function. Atmospheric Chemistry and Physics Discussions, 2004, 4 (1), pp.39-69. hal-00300852

HAL Id: hal-00300852

<https://hal.science/hal-00300852>

Submitted on 6 Jan 2004

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**A new time-resolved
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Received: 17 October 2003 – Accepted: 4 December 2003 – Published: 6 January 2004

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Abstract

A time-resolved model of the Na layer in the mesosphere/lower thermosphere region is described, where the continuity equations for the major sodium species Na, Na⁺ and NaHCO₃ are solved explicitly, and the other short-lived species are treated in steady-state. It is shown that the diurnal variation of the Na layer can only be modelled satisfactorily if sodium species are permanently removed below about 85 km, both through the dimerization of NaHCO₃ and the uptake of sodium species on meteoric smoke particles that are assumed to have formed from the recondensation of vaporized meteoroids. When the sensitivity of the Na layer to the meteoroid input function is considered, an inconsistent picture emerges. The ratio of the column abundances of Na⁺ to Na is shown to increase strongly with the average meteoroid velocity, because the Na is injected at higher altitudes. Comparison with a limited set of Na⁺ measurements indicates that the average meteoroid velocity is probably less than about 25 km s⁻¹, in agreement with velocity estimates from conventional meteor radars, and considerably slower than recent observations made by wide aperture incoherent scatter radars. The Na column abundance is shown to be very sensitive to the meteoroid mass input rate, and to the rate of vertical transport by eddy diffusion. Although the magnitude of the eddy diffusion coefficient in the 80–90 km region is uncertain, there is a consensus between recent models using parameterisations of gravity wave momentum deposition that the average value is less than $3 \times 10^5 \text{ cm}^2 \text{ s}^{-1}$. This requires that the global meteoric mass input rate is less than about 20 t d⁻¹, which is closest to estimates from incoherent scatter radar observations. Finally, the diurnal variation in the meteoroid input rate only slight perturbs the Na layer, because the residence time of Na in the layer is several days, and diurnal effects are effectively averaged out.

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1. Introduction

The first objective in this paper is to describe a new time-resolved model of the sodium layer that occurs in the upper mesosphere/lower thermosphere (MLT). We have developed this model, hereafter termed NAMOD, to provide a more realistic description of the input of Na atoms into the MLT from meteoric ablation, the conversion of Na atoms into Na⁺ ions and small Na-containing molecules, and the eventual removal of these species from the gas phase via condensation onto particles. The model also permits the behaviour of the Na layer to be analysed on a much shorter time-scale than is possible using traditional steady-state Na models (Plane, 1991; McNeil et al., 1995; Plane et al., 1999). The second objective is to address the large degree of uncertainty, illustrated in this special issue of Atmospheric Chemistry and Physics, concerning the meteoroid velocity distribution and the total meteoroid mass entering the atmosphere. The purpose here is to investigate whether modelling the characteristic features of the Na layer constrains the widely ranging estimates of these meteoroid input parameters.

Figure 1 is a schematic diagram of the gas-phase chemistry of sodium that is believed to take place in the MLT (McNeil et al., 2002). Almost all of these reactions have now been studied in isolation in the laboratory, under conditions appropriate to the MLT (Plane, 2002). Above the atomic Na layer around 90 km, ion chemistry predominates. Na atoms are ionized mostly by charge transfer with the ambient NO⁺ and O₂⁺ ions, together with a small contribution from solar photo-ionization. Na⁺ is then neutralised by forming clusters which undergo dissociative electron recombination. Below 90 km, Na becomes converted to the stable reservoir NaHCO₃, via a series of steps beginning with the oxidation of Na by O₃ to NaO. As shown in Fig. 1, species such as NaO, NaOH and NaHCO₃ are converted back to Na by reaction with O and H, and also by photolysis.

The most common type of Na layer model that has been employed is where the metal chemistry is assumed to be in steady state at each height level in the model (Plane, 1991; McNeil et al., 1995; Plane et al., 1999; McNeil et al., 2002). Such

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a model assumes that the gas-phase chemistry is fast on the time-scale of vertical transport, and that the metal chemistry is closed. That is, all the constituent species cycle between each other, so that no species provides a permanent sink in the upper mesosphere. However, in a recent laboratory study of the photochemistry of sodium species (Self and Plane, 2002), we have found that the major reservoir, NaHCO_3 , has a sufficiently large photolysis cross-section below 240 nm that its photolysis lifetime is only a few hours in the MLT. As will be demonstrated in Sect. 2.4, this would cause a large increase in Na below 90 km during the day, arising from NaHCO_3 photolysis. Since such an increase is not in fact observed (Clemesha et al., 1982; States and Gardner, 2000), the most likely explanation is that NaHCO_3 is permanently removed in the 75 to 85 km region. In the new model to be described here, this removal occurs through the polymerization of NaHCO_3 into dimers, and also by the uptake of sodium species onto meteoric smoke particles (Hunten et al., 1980; Kalashnikova et al., 2000). Hence, a steady-state treatment of the chemistry that partitions sodium between Na and NaHCO_3 is no longer appropriate.

The major source of sodium in the MLT is the ablation of sporadic meteoroids, which provide a much greater mass flux than meteor showers (Ceplecha et al., 1998; Baggaley, 2002). Unfortunately, the average daily input of meteoroids into the atmosphere is rather uncertain. Until the 1990s, estimates of the mass flux passing through the beams of meteor radars indicated a figure of about 44 t d^{-1} over the entire Earth (Hughes, 1978). However, conventional meteor radars actually measure specular reflections from the ion trail created in the atmosphere by the ablating meteoroid, from which the mass and velocity of the meteoroid are derived indirectly (Baggaley, 2002). Furthermore, the wavelength of the radar will only sample a subset of the mass/velocity distribution of the meteoroids. In fact, the mass flux from conventional radar estimates has now been revised upwards by a factor of 2–3 (Hughes, 1997).

The meteoroid mass input rate has also been estimated by other techniques. For instance, measuring the accumulation of iridium in ocean-floor sediments indicates that the meteoric influx could be up to 240 t d^{-1} (Wasson and Kyte, 1987). More recently,

the analysis of small particle impact craters on the Long Duration Exposure Facility (LDEF), an orbital impact detector placed on a spacecraft for several years, has yielded an estimate of $80\text{--}130\text{ t d}^{-1}$ (Love and Brownlee, 1993; McBride et al., 1999). The median mass of the incoming dust particles was deduced to be $10\text{ }\mu\text{g}$, with a median radius of $100\text{ }\mu\text{m}$ and a mean entry velocity of about 18 km s^{-1} . Single-particle analyses of stratospheric aerosol have shown that half of the particles in the lower stratosphere contain 0.5 to 1.0% meteoric iron by mass, requiring a total extraterrestrial flux of between 20 and 100 t d^{-1} (Cziczo et al., 2001). In the past couple of years, high-powered large aperture radars, such as the radio telescope at Arecibo, have reported direct observations of the meteor head echo (i.e. the ball of plasma around the ablating particle as it descends through the atmosphere). This enables measurements of the direction of origin, velocity, deceleration and mass to be made (Pellinen-Wannberg and Wannberg, 1994; Mathews et al., 1997; Close et al., 2000; Janches et al., 2000; Mathews et al., 2001). The observed mass distribution is shifted to smaller mass ranges (median mass $\approx 1\text{ }\mu\text{g}$) compared with the LDEF results, so that the total incoming mass rate is about 10 t d^{-1} ; that is, 1 order of magnitude lower than the LDEF estimate. The mean entry velocity is also significantly higher, around $40\text{--}50\text{ km s}^{-1}$. These conflicting results from the two different types of radar measurement have yet to be resolved.

2. The time-resolved sodium model (NAMOD)

2.1. Sodium chemistry

The sodium reaction scheme used in NAMOD is illustrated in Fig. 1. The appropriate rate coefficients, listed in Table 1, are taken from a recent review by the author (Plane, 2002). Note that a number of minor sodium species (e.g. NaO_3 and NaCO_3) and minor reaction pathways have been omitted from the model in the interests of computing efficiency. This has a negligible effect on the computed atomic Na profiles.

The rate constant for reaction 12 in Table 1, the dimerization of NaHCO_3 to form

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(NaHCO₃)₂, was calculated using Rice-Ramsberger-Kassell-Markus (RRKM) theory, which we have recently applied to recombination reactions of metallic species (Plane and Rollason, 2001). The RRKM calculation requires as input the molecular parameters (optimized geometries, moments of inertia, vibrational frequencies and dipole moments) of NaHCO₃ and (NaHCO₃)₂. These were calculated using the hybrid density functional/Hartree-Fock B3LYP method from within the Gaussian 98 suite of programs (Frisch et al., 1998), combined with the 6-311+g(2d, p) triple zeta basis set. This is a large, flexible basis set which has both polarization and diffuse functions added to the atoms (Foresman and Frisch, 1996). Fig. 2 shows the optimized geometry of (NaHCO₃)₂. There is strong electrostatic binding in the dimer because NaHCO₃ has a substantial dipole moment (6.8 Debye). The resulting binding energy of the dimer is 180 kJ mol⁻¹ (with an expected uncertainty of ±4 kJ mol⁻¹ at this level of theory, Foresman and Frisch, 1996). Moreover, the dimer contains six vibrational modes with very low frequencies (<100 cm⁻¹), which are really free internal rotors. The resulting density of rovibrational states is so high that the recombination reaction to form (NaHCO₃)₂ is essentially at the high pressure limit, even in the MLT pressure range of 10⁻⁴–10⁻³ mbar. Hence, *k*₁₂ is given by long-range dipole-dipole capture theory (Stoecklin and Clary, 1992):

$$k_{12} = 1.766 \times 10^6 \left(\frac{\pi}{M_r} \right)^{1/2} \left(\frac{\mu_{\text{NaHCO}_3}^2}{4\pi\epsilon_0} \right)^{2/3} (k_B T)^{-1/6}, \quad (1)$$

where *M_r* is the reduced mass of two NaHCO₃ molecules (=6.98×10⁻²⁶ kg molecule⁻¹), *μ*_{NaHCO₃} is the dipole moment of NaHCO₃ (=2.27×10⁻²⁹ cm), *ε*₀ is the permittivity in a vacuum, and *k_B* is the Boltzmann constant. Note that according to these quantum calculations the dimer should be a sink for sodium, because it will not photolyse at wavelengths above 180 nm, nor react with atomic H (in contrast to NaHCO₃, Cox et al., 2001; Self and Plane, 2002).

In NAMOD, the concentrations of the three major gas-phase sodium species, Na,

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NaHCO_3 and Na^+ , are determined by full solution of their respective continuity equations. In contrast, all the other sodium species shown in Fig. 1 are short-lived intermediates, and are therefore correctly treated in chemical steady state. The continuity equations of the major species are then given by:

$$5 \quad k_{\text{het}}^{\text{Na}} \frac{d[\text{Na}]}{dt} = I_{\text{abl}} + A[\text{NaHCO}_3] + B[\text{Na}^+] - (C + D + k_{\text{het}}^{\text{Na}}) [\text{Na}] - \nabla \Theta^{\text{Na}} \quad (2)$$

$$\frac{d[\text{NaHCO}_3]}{dt} = D[\text{Na}] - (A + k_{\text{het}}^{\text{NaHCO}_3}) [\text{NaHCO}_3] - k_{12}[\text{NaHCO}_3]^2 - \nabla \Theta^{\text{NaHCO}_3} \quad (3)$$

$$\frac{d[\text{Na}^+]}{dt} = C[\text{Na}] - (B[\text{Na}^+] + k_{\text{het}}^{\text{Na}^+}) - \nabla \Theta^{\text{Na}^+}, \quad (4)$$

where I_{abl} is the rate of injection of Na from meteoric ablation (see below),

$$A = (k_9[\text{H}] + J_{33}) [\text{NaHCO}_3],$$

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$$B = \left(k_{22}[\text{N}_2][\text{M}] \left(\frac{k_{24}([\text{CO}_2] + [\text{H}_2\text{O}]) + k_{29}[\text{e}^-]}{k_{24}([\text{CO}_2] + [\text{H}_2\text{O}]) + k_{29}[\text{e}^-] + k_{25}[\text{O}] \frac{(k_{26}[\text{O}] + k_{28}[\text{O}_2])}{(k_{26}[\text{O}] + k_{28}[\text{O}_2] + k_{27}[\text{N}_2])}} \right) + k_{23}[\text{CO}_2][\text{M}] \right)$$

$$C = J_{34} + k_{20}[\text{O}_2^+] + k_{21}[\text{NO}^+],$$

$$D = \left(k_1[\text{O}_3] + k_{10}[\text{O}_2][\text{M}] \frac{k_{11}[\text{O}]}{k_{11}[\text{O}] + J_{31}} \right)$$

$$\left(\frac{k_4[\text{H}_2] + k_6[\text{H}_2\text{O}]}{k_2[\text{O}] + k_3[\text{O}_3] + k_4[\text{H}_2] + k_5[\text{H}_2] + k_6[\text{H}_2\text{O}] + J_{30}} \right) \left(\frac{k_8[\text{CO}_2][\text{M}]}{J_{32} + k_7[\text{H}] + k_8[\text{CO}_2][\text{M}]} \right).$$

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$k_{\text{het}}^{\text{X}}$ is the first-order rate of uptake of species X on meteoric smoke particles (see below), and $\nabla \Phi^{\text{X}}$ is the divergence of the vertical flux of species X.

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The transport of all the sodium constituents is then assumed to be governed by the same eddy diffusion coefficient up to the turbopause at about 100 km, and the vertical flux of species X at height z is given by

$$\Phi^X = -K_{zz} \left(\frac{\partial [X]}{\partial z} + [X] \left(\frac{1}{H} + \frac{1}{T} \frac{\partial T}{\partial z} \right) \right), \quad (5)$$

5 where K_{zz} is the vertical eddy diffusion coefficient, H is the atmospheric scale height, and T the temperature. Above the turbopause, molecular diffusion dominates over eddy diffusion, and K_{zz} and H in Eq. (5) are replaced by D_X and H_X , the molecular diffusion coefficient and the scale height of the individual species X, respectively (Chabrilat et al., 2002). NAMOD is 1 dimensional, extending from 65 to 110 km with a height
 10 resolution of 0.5 km. Equations (2)–(4) are integrated with a 10 min time step using a time-implicit integration scheme (Shimazaki, 1985). The changing diurnal concentration profiles of K_{zz} and the background mesospheric species that control the sodium chemistry (H, O, O₃, NO⁺, e[−] etc.) are read in every 20 model minutes (see below). All model runs shown here were integrated over 20 model days, for the conditions of
 15 January, 40° N.

2.2. Background mesospheric constituents

In order to compute a consistent set of concentration profiles for the background species, a separate mesospheric model (MESOMOD) was run off-line. MESOMOD is also 1 dimensional, extending from 65 to 110 km with a height resolution of 0.5 km.
 20 It contains a full treatment of the odd oxygen and hydrogen chemistry in the MLT, with rate coefficients taken from the NASA/JPL compilation (DeMore et al., 1997). The vertical eddy diffusion coefficient is taken off-line from a 3D global model, TIME-GCM (Kaufmann et al., 2002). This model also supplies the H₂O, H₂ and CO₂ mixing ratios at the lower boundary (65 km). Parameterised codes for the photolysis of O₃, H₂O,
 25 HO₂, H₂O₂, and O₂ at Lyman- α and in the Schumann-Runge continuum, are used (Shimazaki, 1985; Brasseur and Solomon, 1998). The photolysis of O₂ in the Schumann-

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Runge bands is calculated using a recent formalism (Koppers and Murtagh, 1996). MESOMOD uses time-implicit integration with a 2 min time step, and was run for 5 model days before a set of diurnal concentration profiles was recorded every 20 min for input into NAMOD. Note that because the concentration of sodium species is at least 4 orders of magnitude less than species such as O_3 or H in the MLT, the background chemistry is essentially decoupled from the sodium. Profiles of NO^+ , O_2^+ and e^- are input off-line for the appropriate year and time-of-day from the International Reference Ionosphere 2000 (Bilitza, 2003).

2.3. Meteoroid ablation model

The physics of meteoroid ablation, which occurs because the very high entry velocities of meteoroids lead to rapid heating by collision with air molecules, has been treated in detail by several investigators (Hunten et al., 1980; Love and Brownlee, 1991; Kalashnikova et al., 2000; McNeil et al., 2002). The problem becomes reasonably tractable for particles less than about $250 \mu m$ in radius, because heat conductivity through the particle is fast enough for the particle to be treated as isothermal (Love and Brownlee, 1991). Most of the total mass of cosmic dust is made up from particles with masses less than $100 \mu g$ (Love and Brownlee, 1993). Assuming a density of $2 \times 10^3 \text{ kg m}^{-3}$ (typical of an ordinary chondrite), a $250 \mu m$ meteoroid has a mass of $130 \mu g$, thereby justifying the neglect of heat conductivity.

The frictional heating of the meteoroid by collisions with air molecules is balanced by radiative losses, and by the absorption of heat energy through temperature increases, melting, phase transitions, and vaporization. In order to calculate these terms, parameters such as the meteoroid shape, density, and composition are needed (Rietmeijer, 2000, 2002). In NAMOD it is assumed that most of the extra-terrestrial material has the composition of ordinary chondrites. If the particles is spherical (i.e. the shape factor is

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1.2), then the energy balance equation is given by (Jones and Kaiser, 1966):

$$\frac{1}{2}\Lambda\rho_a v^3 = 4\sigma\varepsilon (T_s^4 - T_a^4) + \frac{4}{3}R\rho_m C \frac{dT_m}{dt}. \quad (6)$$

The left-hand side of Eq. (1) represents the frictional heating term, where Λ (= 0.5) is the heat transfer coefficient or fraction of the total kinetic energy of the air molecules that is transferred to the meteoroid, ρ_a is the atmospheric density, and v is the meteoroid velocity. The first term on the right-hand side describes the radiative losses, where σ is Stefan's constant and ε is the emissivity of the meteoroid ($\sigma\varepsilon = 5.7 \times 10^{-8} \text{ J m}^{-2} \text{ s}^{-1}$), T_s is the surface temperature of the meteoroid, and T_a is the ambient atmospheric temperature. The second term represents the energy losses due to heat capacity (i.e. vaporization, phase transitions, and heating); R is the meteoroid radius, ρ_m ($= 2 \times 10^3 \text{ kg m}^{-3}$) the meteoroid density, C ($= 1 \times 10^3 \text{ J K}^{-1} \text{ kg}^{-1}$) the meteoroid specific heat, T_m the mean temperature of the particle, and t is the time. The change in velocity of the meteoroid is given by

$$\frac{dv}{dt} = -\frac{\Gamma\rho_a\pi R^2 v^2}{m} + \frac{\pi}{(z+r_e)^2}. \quad (7)$$

The first term on the right-hand side describes deceleration, where Γ is the atmospheric drag parameter ($= 1.2$) and m is the meteoroid mass. The second term describes gravitational acceleration of the particle, where z is the height in the atmosphere, r_e ($= 6.378 \times 10^6 \text{ m}$) is the radius of the earth, and μ ($= 3.99 \times 10^{14} \text{ m}^3 \text{ s}^{-2}$) is the gravitational constant of the earth. In the present model, all meteoroids enter the atmosphere at 37° to the zenith.

Equation (6) shows that for very small particles, the heat capacity term will be much smaller than the radiative loss term, in which case the meteoroid will not become hot enough to ablate. This is the case for all particles smaller than 100 nm. A $1 \mu\text{m}$ particle must enter the atmosphere at over 40 km s^{-1} in order to heat up to 1200 K, the temperature at which volatile elements such as Na begin to evaporate (Fegley and Cameron,

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1987). All meteoroids larger than $10\text{ }\mu\text{m}$ will reach this ablation temperature. Figure 3a shows the altitude at which a $10\text{ }\mu\text{g}$ particle ($R\approx 100\text{ }\mu\text{m}$) reaches temperatures between 500 and 2500 K, over the range of possible meteoroid entry velocities. Note that particles which originate within the solar system must have entry velocities that range from 11.5 km s^{-1} , for a particle in the same orbit as the Earth, to 72 km s^{-1} for a particle in a retrograde orbit (Baggaley, 2002).

Most of the models developed for estimating the meteoric metal input into the atmosphere (Hunten et al., 1980; Plane, 1991; McNeil et al., 1995; McNeil et al., 2002) assume that the relative metallic abundances in the ablated vapour are given by their meteoritic abundances. The major metallic constituents by weight are then: Mg 12.5%, Fe 11.5%, Al 1.7%, Ni 1.5%, Ca 1.0%, Na 0.6% (Mason, 1971). A more realistic approach involving fractionation has been proposed recently, with the relatively volatile elements such as Na and K evaporating first and refractory elements (such as Ca) evaporating last (McNeil et al., 1998). This differential ablation scheme, which is based on planetary fractionation models (Fegley and Cameron, 1987), is able to explain the large depletion of Ca to Na observed in meteor trails (von Zahn et al., 1999).

The present model contains the following simple treatment of differential ablation. When T_s reaches 1200 K, further frictional heat input to the particle is balanced by vaporization and radiative cooling, and no further temperature rise occurs until the initial 20% of the meteoroid mass has evaporated. This is determined from the relation

$$\xi \frac{dm}{dt} = -\frac{1}{2} \Lambda \rho_a v^3 + 4\sigma\epsilon (T_s^4 - T_a^4), \quad (8)$$

where ξ ($=3\times 10^6\text{ J kg}^{-1}$) is the heat of sublimation of the meteoroid. All of the Na is assumed to evaporate in the first 5% of this fraction. Evaporation then ceases until T_s reaches 2100 K, when 78% of the mass evaporates. The final 2%, containing most the Ca and other refractory elements, then evaporates if T_s reaches 2500 K.

Figure 3b shows the mass ablation rate as a function of altitude, for the meteoroid mass and velocity distribution estimated from the LDEF experiment (McBride et al., 1999). This will be termed the LDEF profile. The peak rate of ablation occurs at 91 km,

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and about 62% of the total incoming material vaporizes. The unablated fraction consists of particles that are too small, or that enter the atmosphere too slowly, to vaporize completely. This ablation profile is similar in peak height and width to that observed by conventional meteor radars, such as in a recent paper using the 53.5 MHz radar at K hlungsborn, a mid-latitude site (Singer et al., 2003). Also shown in Fig. 3b is the ablation profile predicted if the velocity distribution is shifted so that the mean velocity increases from 18 to 40 km s⁻¹, in line with the large aperture radar measurements (Pellinen-Wannberg and Wannberg, 1994; Mathews et al., 1997; Close et al., 2000; Janches et al., 2000). This ablation profile, which will be referred to as the ISR profile, now peaks just above 100 km, and 99% of the material ablates. The profile is in sensible accord with a very recent micrometeor profile from the Arecibo Observatory (Mathews et al., 2003), where the observed broad velocity distribution between 40 and 60 km s⁻¹ results in the profile peaking even higher in the atmosphere at 106 km. In the NAMOD model, these profiles are then modulated with a diurnal and seasonal variation taken from a recent radio meteor survey (Yrjola and Jenniskens, 1998).

2.4. Removal of sodium species from the gas phase

Figure 4a illustrates the modelled Na layer at midday and midnight, when no permanent removal of sodium species is included in NAMOD i.e. k_{12} and k_{het}^X are set to zero. This shows that the Na layer always peaks below 90 km, that the peak moves 2 km lower in daytime, and that the underside of the layer has a scale height less than 1 km. These model predictions are not in accord with lidar observations of the diurnal behaviour of the layer (Clemesha et al., 1982; States and Gardner, 2000). However, if removal of NaHCO₃ by dimerization is now switched on in the model, Fig. 4b shows that the layer now peaks around 90 km with little diurnal variation in the peak height, and the layer has an approximately Gaussian shape, in much better agreement with observations (Plane et al., 1999; States and Gardner, 2000).

Figure 4c illustrates the effect of also switching on the heterogeneous removal of the three major sodium species on meteoric smoke particles. These particles have been

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proposed to form through the recondensation of vaporized meteoroids (Hunten et al., 1980; Kalashnikova et al., 2000). The height profile of 1.3 nm radius smoke particles, illustrated in Fig. 4c, is calculated in NAMOD from the total ablated meteoroid flux (Hunten et al., 1980). Note that these small particles, which are really macromolecules (e.g. a cluster of 60 FeSiO₃ molecules would have a radius of about 1.3 nm), cannot be much larger or their sedimentation velocity would increase to the point where they could not remain at significant concentrations above 80 km (Hunten et al., 1980).

The uptake of sodium species on these smoke particles is likely to be highly efficient, and an uptake coefficient of 0.5 is assumed in NAMOD. In fact, NaHCO₃ and Na⁺ should adsorb readily because of electrostatic effects, and we have recently shown in the laboratory that Na atoms stick with high efficiency on a silica surface (B. J. Murray and J. M. C. Plane, University of East Anglia). As shown in Fig. 4c, there is now less than 0.5 km change in the peak height between midnight and midday. In contrast, the centroid height of the layer decreases from 92.5 km to 91.0 km in daytime, which is in good accord with observations (Clemesha et al., 1982). This decrease arises largely because during daytime there is additional ionization of Na on the topside, as well as photolysis of NaHCO₃ on the bottomside. The Na concentration of around 5100 atom cm⁻³ at the layer peak, and the column density of 5.9×10⁹ atom cm⁻², are typical values during winter at mid-latitudes (Plane, 1991).

Note that in these model runs the LDEF velocity distribution was used with a diurnal average Na input flux from ablation set to 3.8×10³ atom cm⁻² s⁻¹, equivalent to a global meteoroid mass input rate of 12.1 t d⁻¹. This Na input flux, together with permanent removal of sodium by heterogeneous uptake and dimerization of NaHCO₃, is the base case in the sensitivity tests that follow.

3. Sensitivity tests on the meteor input function

In this section the NAMOD model is used to investigate which features of the Na layer are particularly sensitive to the meteoroid entry velocity and mass flux. The sensitivity

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of the layer to diurnal variations in the meteoroid flux will also be considered.

3.1. Entry velocity

Figure 5a shows the profiles of the Na injection rate, I_{abl} , predicted for the LDEF and ISR velocity distributions. These profiles peak at altitudes of 94 and 112 km, respectively, a substantial separation of 18 km. Note that these injection profiles peak at altitudes 3 and 10 km higher than the respective mass ablation rate profiles in Fig. 3b, because all the Na is assumed to ablate in the first 5% of vaporized mass. Figure 5b illustrates the predicted profiles of meteoric smoke particles for the two velocity distributions. Because the sedimentation velocity of 1.3 nm particles becomes significant above 90 km ($>2 \text{ cm s}^{-1}$), the profiles both exhibit a sharp cut-off between 85 and 90 km. More particles are produced in the ISR profile because essentially all the incoming meteoroid mass ablates, compared with only 62% for the LDEF profile.

Figure 5c illustrates the predicted diurnally-averaged height profiles of Na, Na^+ and NaHCO_3 for the two velocity distributions. One effect of the greater number of smoke particles in the ISR case is to reduce both NaHCO_3 and Na by increased heterogeneous uptake below 90 km, with the result that the peak of the Na layer moves up from 91 to 93 km. However, although a peak of 93 km is about 2 km higher than typically observed in mid-latitude winter conditions (Plane et al., 1999), this is probably not of significance because of uncertainties in the smoke profile and uptake coefficients of the sodium species. Of greater interest is the ratio of the column abundances of Na^+ to Na, which increases from 0.27 in the LDEF case to greater than 0.5 for the ISR velocity distribution. This occurs because in the ISR case most of the Na is injected well above 100 km (Fig. 5a), where it is converted to Na^+ (by charge transfer with NO^+ and O_2^+) more rapidly than being transported downwards by eddy diffusion.

Although there is only a limited set of rocket-borne mass spectrometric measurements of Na^+ ions in the MLT (Kopp, 1997), we have averaged the data from nine flights at mid- and high latitudes to obtain an approximate Na^+ column abundance of $1.1 \times 10^9 \text{ cm}^{-2}$. That is, the Na^+/Na ratio should be around 0.2–0.25. Comparison with

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the NAMOD predictions above implies that an average entry velocity around 40 km s^{-1} is significantly too high. In order to produce a Na^+/Na ratio below 0.3, most Na must be injected below 100 km, requiring that the average entry velocity be less than about 25 km s^{-1} .

It should be noted that lidar observations of Na in freshly-ablated meteor trails recorded a height distribution with a centroid height of $89.0 \pm 0.3 \text{ km}$, and a root-mean-square width of $3.3 \pm 0.2 \text{ km}$ (Kane and Gardner, 1993). In comparison, the Na ablation profile from the LDEF mass/velocity distribution (Fig. 5a) is both higher (centroid = 96.5 km) and broader (rms width = 8.3 km). However, as Kane and Gardner pointed out, the lidar observations are biased towards larger particles that produce observable concentrations of Na above the background Na layer. Furthermore, since the rate of molecular diffusion varies inversely with the pressure, trails produced higher up are increasingly likely to have diffused into the background.

3.2. Total meteoroid mass

In Sect. 1 there is a discussion of the widely differing estimates of the total meteoroid mass input to the atmosphere, which range from less than 10 to over 200 t d^{-1} . The Na column abundance is clearly sensitive to the meteoroid flux, and so it might be thought that application of a model such as NAMOD would provide a means of constraining these flux estimates. However, the Na column abundance is also controlled by the chemistry in the MLT, and the rate of downward transport by eddy diffusion. While laboratory experiments appear to have placed the gas-phase chemistry on reasonably firm ground (Plane, 2002), there is currently uncertainty regarding the appropriate values of K_{zz} to use for describing vertical transport in the MLT. Atmospheric models which include a parameterisation of gravity wave momentum deposition predict an average value for K_{zz} between 80 and 90 km of $(1-4) \times 10^5 \text{ cm}^2 \text{ s}^{-1}$ (Garcia and Solomon, 1994; Chabrilat et al., 2002; Kaufmann et al., 2002). In contrast, estimates based on measured energy dissipation rates in the MLT find values up to an order of magnitude

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greater (Hocking, 1990).

Figure 6a shows the effect on the diurnally-averaged Na layer of increasing K_{zz} by factors of 2, 5, 10 and 20, for a constant meteoric mass input (12.1 t d^{-1}). Note the marked reduction in the column abundance and peak density, with increasing K_{zz} . The peak height of the layer also decreases from 91.0 to 88.5 km. However, this change is comparatively modest, and because the peak height is also sensitive to uncertainties in the chemistry, it is not as useful a test of the balance between the meteoric input flux and downward transport by eddy diffusion.

Figure 6b is a contour plot showing the Na column abundance predicted by NAMOD for a range of meteoroid mass inputs (1.4 to 170 t d^{-1}) and values of K_{zz} (8×10^4 to $4.6 \times 10^6 \text{ cm}^2 \text{ s}^{-1}$) between 80 and 90 km. This plot demonstrates the sensitivity of the Na column abundance to both parameters. However, inspection of Fig. 6b shows that the typical Na column abundance of $5 \times 10^9 \text{ cm}^{-2}$ can be produced from a range of combinations of these parameters. For instance, this column abundance is predicted to result from a meteoroid mass input around 12 t d^{-1} , as estimated by the ISR observations (Sect. 2.3), combined with a low K_{zz} value around $2 \times 10^5 \text{ cm}^2 \text{ s}^{-1}$, as predicted by gravity wave momentum deposition models (Kaufmann et al., 2002). But the same column abundance can be produced by the higher mass input estimated from conventional meteor radars ($40\text{--}50 \text{ t d}^{-1}$) (Hughes, 1978), if a larger value of K_{zz} around $10^6 \text{ cm}^2 \text{ s}^{-1}$ is adopted, consistent with measured energy dissipation rates (Hocking, 1990). Finally, mass input rates above 100 t d^{-1} , estimated from the LDEF experiment (Love and Brownlee, 1993) and the accumulation of iridium in deep-sea sediments (Wasson and Kyte, 1987), require very large values of K_{zz} ($>3 \times 10^6 \text{ cm}^2 \text{ s}^{-1}$) in order to limit the build up of Na in the MLT.

Since such large values of the eddy coefficient are greater than both experimental (Hocking, 1990) and theoretical (Kaufmann et al., 2002) estimates of K_{zz} , it seems that the global mass input rate must be below 100 t d^{-1} . Furthermore, a recent modelling study of the significant decrease in the CO_2 mixing ratio above 80 km has demonstrated that the average K_{zz} between 80 and 90 km needs to be less than about $3 \times 10^5 \text{ cm}^2 \text{ s}^{-1}$

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in order for molecular diffusion of CO₂ to compete with eddy diffusion, and hence produce some gravitational separation (Chabrilat et al., 2002). Figure 6b shows that this upper limit to K_{zz} corresponds to an upper limit to the global mass input rate of about 20 t d⁻¹.

Of course, one uncertainty in this analysis is the assumption in the ablation model that all the Na ablates in the first 5% of mass lost once the meteoroid reaches a temperature of 1200 K (Sect. 2.3). In fact, there is essentially no change in the Na ablation flux when the “boiling point” for Na ablation is raised to 2000 K. This is because only a small subset of meteoroids that reach 1200 K do not continue heating to 2000 K. The assumption that the Na ablates in the first mass fraction is based on a thermodynamic model of high temperature melts (Fegley and Cameron, 1987), and diffusion of Na through the bulk of a meteoroid should not constrain the rate of evaporation from a molten particle smaller than about 200 μm. There is also some uncertainty in various parameters in the ablation model, including the meteoroid density (ρ_m), the heat capacity (C), and the heat of sublimation (ξ), whose values in the model are based on those of chondritic meteorites (Hunten et al., 1980). However, if all Na ablates in the initial 5% of mass lost, then the ablation flux of Na is not in fact sensitive to these parameters within reasonable limits (factor of 2 variation), and changing these parameters only affects the height profile of I_{abl} .

3.3. The diurnal variation in the meteoroid input

Finally, we consider the sensitivity of the Na layer to the diurnal variation in meteoroid flux. In the case of radio meteors, there is a dawn peak and late afternoon minimum, in the ratio of 5:1 (Yrjola and Jenniskens, 1998). ISR observations suggest an even more sharply peaked distribution around dawn (Janches et al., 2001; Mathews et al., 2001). In order to examine the sensitivity of the Na layer to the diurnal variation, NAMOD was run for two extreme cases where the integrated diurnal input was the same: 1), all the input occurs between 05:00 and 07:00 local time; and 2), the input is constant over the diurnal cycle. Figure 7 is a contour plot of the difference in the Na concentration

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between cases 1 and 2, as a function of height and local time. While this appears to show a dramatic difference around dawn, it should be borne in mind that the peak Na concentration around 90 km is over 5000 cm^{-3} (Fig. 4c), and so the largest difference in Na concentration shown in Fig. 7 is actually only 1.5% of the concentration in the layer. This change is small because the residence time of Na in the layer is about 6 days, so that diurnal variations in the meteoroid input function are essentially averaged out.

4. Conclusions

A new time-resolved model of the Na layer (NAMOD) has been described. The model runs efficiently by solving the continuity equations for the major sodium species Na, Na^+ and NaHCO_3 explicitly, and placing the other short-lived species in steady-state. It is shown that the diurnal variation of the Na layer can be modelled satisfactorily only if sodium is permanently removed, both through dimerization of NaHCO_3 and the uptake of sodium species on meteoric smoke particles that have formed from the recondensation of vaporized meteoroids.

When the sensitivity of the Na layer to the meteoroid input function is considered, a rather inconsistent picture emerges. The ratio of the column abundances of Na^+ to Na increases with the average meteoroid velocity: comparison with a limited set of Na^+ measurements indicates that the average meteoroid velocity is probably below 25 km s^{-1} , as observed by conventional meteor radars. The Na column abundance is sensitive to both the meteoroid mass input rate and the vertical eddy diffusion coefficient controlling downward transport. Although there is uncertainty regarding K_{zz} , there is a consensus from models of gravity wave momentum deposition, and modelling the observed fall-off in the CO_2 mixing ratio in the MLT, that the upper limit to K_{zz} in the 80 to 90 km region is about $3 \times 10^5 \text{ cm}^2 \text{ s}^{-1}$. This implies that the mass input rate is less than 20 t d^{-1} , close to the estimates from wide aperture incoherent scatter radars. Finally, the diurnal variation in the meteoroid input rate causes only a slight

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perturbation to the Na layer, because the residence time of Na in the layer is several days and diurnal effects tend to be averaged out.

Acknowledgements. This work was supported by the UK Natural Environment Research Council and the US Air Force Research Laboratory. I am grateful to R. G. Roble (High Altitude Observatory, Boulder, Colorado) for supplying output from the TIME-GCM model.

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Table 1. Gas-phase neutral and ion-molecule reactions of Na Species contained in the model.

Number	Reaction	Rate coefficient [#]
Neutral chemistry		
(R1)	$\text{Na} + \text{O}_3 \rightarrow \text{NaO} + \text{O}_2$	$1.1 \times 10^{-9} \exp(-116/T)$
(R2)	$\text{NaO} + \text{O} \rightarrow \text{Na} + \text{O}_2$	$2.2 \times 10^{-10} (T/200)^{1/2}$
(R3)	$\text{NaO} + \text{O}_3 \rightarrow \text{Na} + 2\text{O}_2$	$3.2 \times 10^{-10} \exp(-550/T)$
(R4)	$\text{NaO} + \text{H}_2 \rightarrow \text{NaOH} + \text{H}$	$1.1 \times 10^{-9} \exp(-1100/T)$
(R5)	$\text{NaO} + \text{H}_2 \rightarrow \text{Na} + \text{H}_2\text{O}$	$1.1 \times 10^{-9} \exp(-1400/T)$
(R6)	$\text{NaO} + \text{H}_2\text{O} \rightarrow \text{NaOH} + \text{OH}$	$4.4 \times 10^{-10} \exp(-507/T)$
(R7)	$\text{NaOH} + \text{H} \rightarrow \text{Na} + \text{H}_2\text{O}$	$4 \times 10^{-11} \exp(-550/T)$
(R8)	$\text{NaOH} + \text{CO}_2 (+\text{M}) \rightarrow \text{NaHCO}_3$	$1.9 \times 10^{-28} (T/200)^{-1}$
(R9)	$\text{NaHCO}_3 + \text{H} \rightarrow \text{Na} + \text{H}_2\text{CO}_3$	$1.84 \times 10^{-13} T^{0.777} \exp(-1014/T)$
(R10)	$\text{Na} + \text{O}_2 (+\text{M}) \rightarrow \text{NaO}_2$	$5.0 \times 10^{-30} (T/200)^{-1.22}$
(R11)	$\text{NaO}_2 + \text{O} \rightarrow \text{NaO} + \text{O}_2$	$5.0 \times 10^{-10} \exp(-940/T)$
(R12)	$2\text{NaHCO}_3 (+\text{M}) \rightarrow (\text{NaHCO}_3)_2$	$8.8 \times 10^{-10} (T/200 \text{ K})^{-0.23}$
Ion-molecule chemistry		
(R20)	$\text{Na} + \text{O}_2^+ \rightarrow \text{Na}^+ + \text{O}_2$	2.7×10^{-9}
(R21)	$\text{Na} + \text{NO}^+ \rightarrow \text{Na}^+ + \text{NO}$	8.0×10^{-10}
(R22)	$\text{Na}^+ + \text{N}_2 (+\text{M}) \rightarrow \text{Na.N}_2^+$	$4.8 \times 10^{-30} (T/200)^{-2.2}$
(R23)	$\text{Na}^+ + \text{CO}_2 (+\text{M}) \rightarrow \text{Na.CO} + 2^+$	$3.7 \times 10^{-29} (T/200)^{-2.9}$
(R24)	$\text{Na.N}_2^+ + \text{X} \rightarrow \text{Na.X}^+ + \text{N}_2 (\text{X}=\text{CO}_2, \text{H}_2\text{O})$	6×10^{-10}
(R25)	$\text{Na.N}_2^+ + \text{O} \rightarrow \text{NaO}^+ + \text{N}_2$	4×10^{-10}
(R26)	$\text{NaO}^+ + \text{O} \rightarrow \text{Na}^+ + \text{O}_2$	1×10^{-11}
(R27)	$\text{Na.O}^+ + \text{N}_2 \rightarrow \text{Na.N}_2^+ + \text{O}$	1×10^{-12}
(R28)	$\text{Na.O}^+ + \text{O}_2 \rightarrow \text{Na}^+ + \text{O}_3$	5×10^{-12}
(R29)	$\text{Na.Y}^+ + \text{e}^- \rightarrow \text{Na} + \text{Y} (\text{Y}=\text{N}_2, \text{CO}_2, \text{H}_2\text{O}, \text{O})$	$1 \times 10^{-6} (T/200)^{-1/2}$

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Table 1. Continued.

Number	Reaction	Rate coefficient [#]
	Photochemical reactions	
(R30)	$\text{NaO} + h\nu \rightarrow \text{Na} + \text{O}$	5.5×10^{-2}
(R31)	$\text{NaO}_2 + h\nu \rightarrow \text{Na} + \text{O}_2$	1.9×10^{-2}
(R32)	$\text{NaOH} + h\nu \rightarrow \text{Na} + \text{OH}$	1.8×10^{-2}
(R33)	$\text{NaHCO}_3 + h\nu \rightarrow \text{Na} + \text{HCO}_3$	1.3×10^{-4}
(R34)	$\text{Na} + h\nu \rightarrow \text{Na}^+ + \text{e}^-$	2×10^{-5}

[#] Rate coefficients are taken from a recent review of mesospheric metal chemistry (Plane, 2002), except for R12 (see text).

Units: unimolecular, s^{-1} ; bimolecular, $\text{cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$; termolecular, $\text{cm}^6 \text{ molecule}^{-2} \text{ s}^{-1}$

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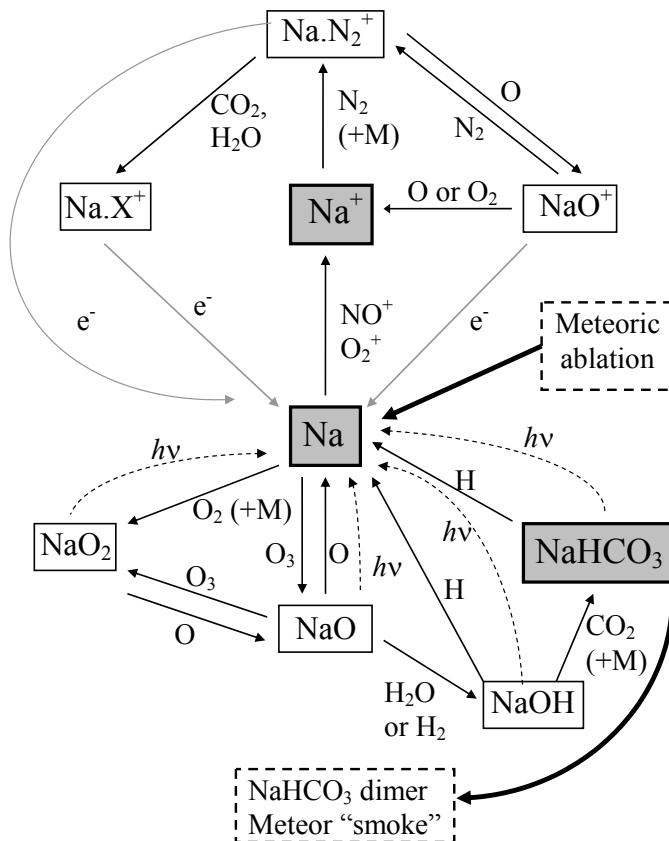


Fig. 1. Schematic drawing of the sodium chemistry in the NAMOD model. The thin solid lines indicate reactions with measured rate coefficients. The broken lines are measured photodissociation reactions. The grey solid lines are dissociative electron recombination reactions with estimated rate coefficients.

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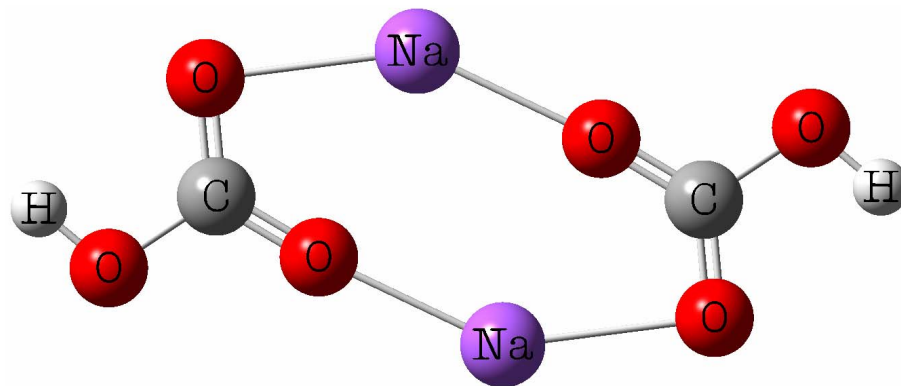


Fig. 2. The sodium bicarbonate dimer, calculated at the B3LYP/6-311+G(2d, p) level of theory. The Na atoms are 3.2 Å apart, and the H atoms are 8.9 Å apart.

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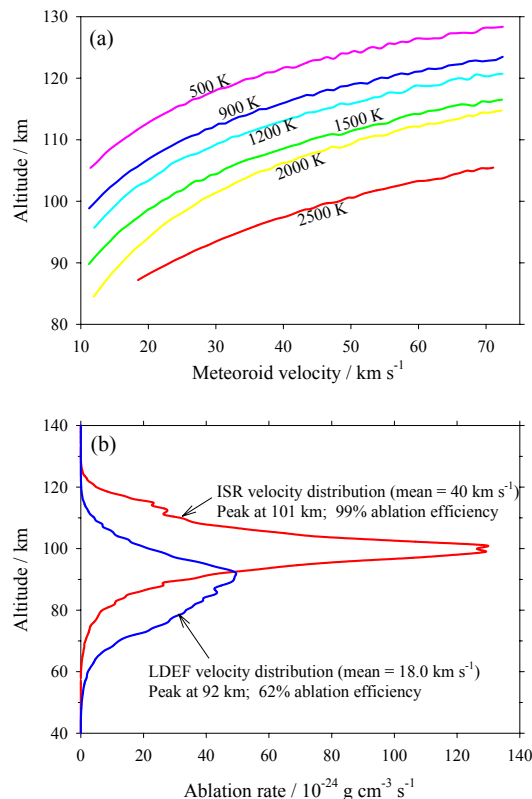


Fig. 3. (a) The height at which a 10 μg meteoroid reaches the temperature shown on each of the curves, as a function of meteoroid entry velocity. (b) Height profiles of the total mass ablation rate for the LDEF meteoroid size distribution (reduced by 85% in all mass ranges to give a global input of 12.1 t d⁻¹), calculated using the assumed LDEF velocity distribution (blue line), or the LDEF velocity distribution shifted by 22 km s⁻¹ to simulate the velocity distribution observed by incoherent scatter radars (red line).

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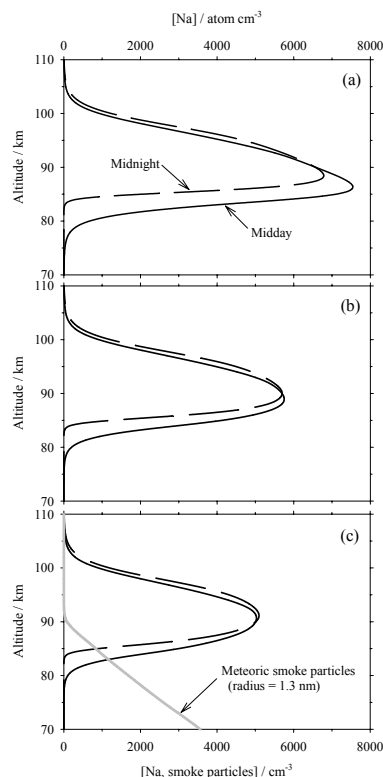


Fig. 4. Modelled height profiles of the Na layer at midday and midnight, for 40° N, January: **(a)**, no permanent sinks for gas-phase sodium species; **(b)**, dimerization of NaHCO_3 is included; **(c)**, dimerization of NaHCO_3 and permanent removal of sodium species on meteoric smoke particles of radius 1.3 nm are included. The smoke particle profile is depicted by the solid grey line. The diurnal average Na ablation flux is $3.8 \times 10^3 \text{ atom cm}^{-2} \text{ s}^{-1}$.

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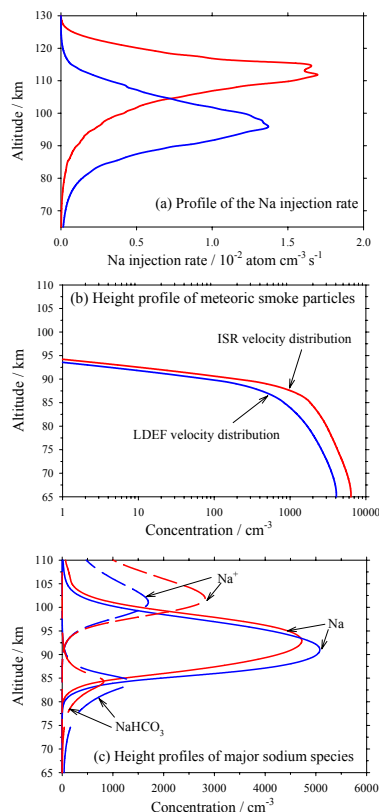


Fig. 5. (a) Height profiles of the Na input rate from meteor ablation, calculated for the LDEF meteoroid size distribution (reduced by 85% in all mass ranges to give a global input of 12.1 t d^{-1}) and the LDEF velocity distribution (blue line), or the LDEF velocity distribution shifted by 22 km s^{-1} to simulate the velocity distribution observed by incoherent scatter radars (red line). (b) Height profiles of the concentration of 1.3 nm radius meteoric smoke particles, for the two velocity distributions in (a). (c) Diurnally-averaged height profiles of Na , Na^+ and NaHCO_3 , for the two velocity distributions in (a). Conditions are 40°N , January.

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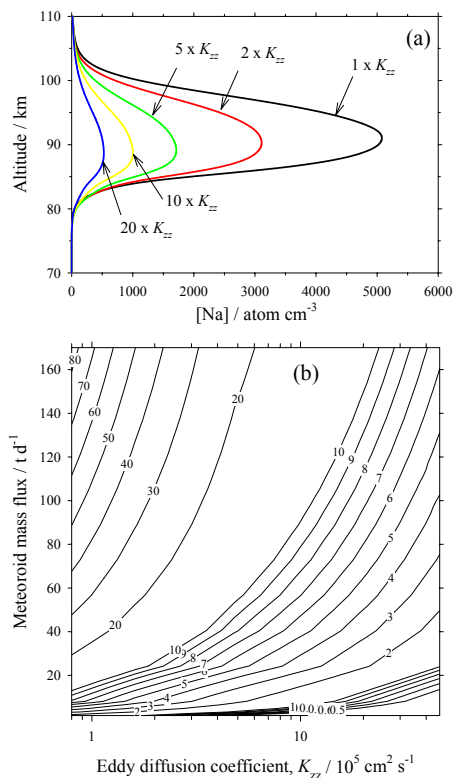


Fig. 6. (a) Diurnally-averaged Na profiles where the eddy diffusion coefficient profile is multiplied by the factor indicated. The global meteoroid mass input rate is 12.1 t d⁻¹. **(b)** Contour plot of the diurnally-averaged Na column density (in 10⁹ atom cm⁻²) as a function of the global meteoroid mass input rate (assuming the LDEF mass and velocity distribution), and the average eddy diffusion coefficient between 80 and 90 km. The typical observed column density is about 5 x 10⁹ atom cm⁻². Conditions: 40° N, January.

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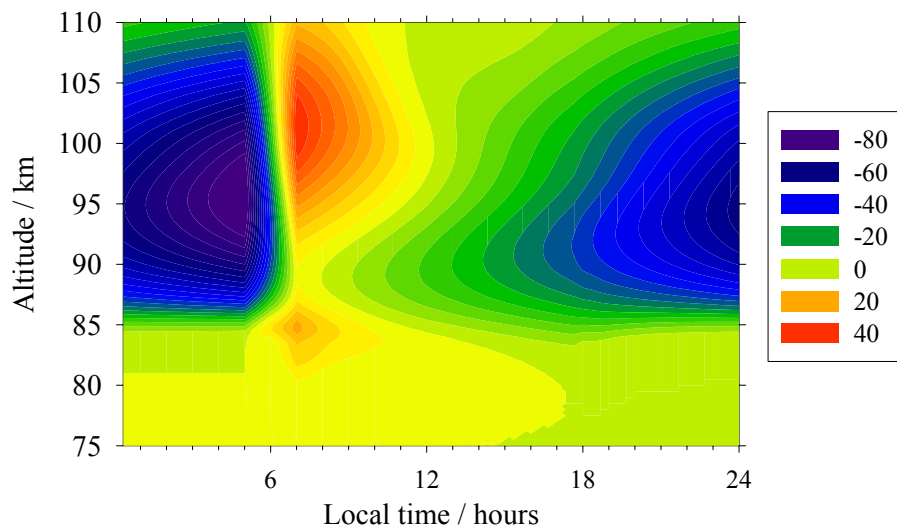


Fig. 7. Contour plot showing the difference between the Na atom density (atom cm^{-3}), as a function of height and local time, for the cases where all meteoric input occurs between 05:00 and 07:00 LT, and where the meteoric input is constant over a diurnal cycle.

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